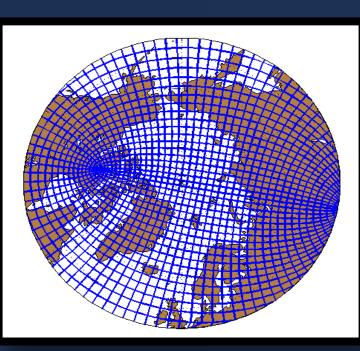
Hindcast Simulations of Arctic and Sub-Arctic Oceans S. Lundrigan (sarah.lundrigan@mun.ca), E. Demirov (entcho@mun.ca) Memorial University of Newfoundland

MEMORIAL UNIVERSITY

Model and Experimental design

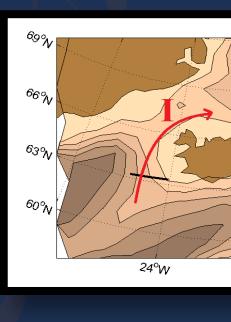
oThe ocean model NEMO-OPA and the sea-ice model NEMO-LIM are configured on a **2**° grid with 31 vertical levels.



olt is forced using 6 hourly atmospheric conditions from NCEP/NCAR reanalysis for the period 1948-2005.

OThe results shown here are from six ensemble runs initialized from the control run model solution for years with high, low and neutral NAO index.

Greenland-Scotland Ridge model results than in the observations. oThe salinity from are within error of salinity.

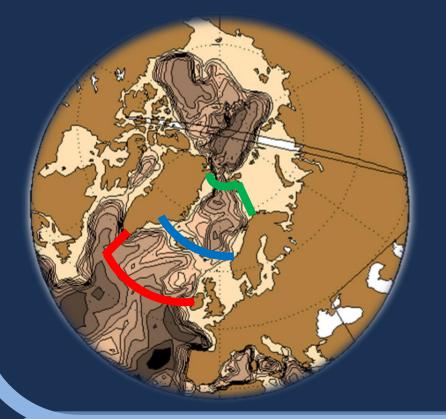


Interannual Variability

• There was an increase in the decadal mean of the temperature during the studied period in the Barents Sea Opening and Fram Strait while the temperature over the Greenland-Scotland Ridge showed no long term trend.

OThe decadal mean of volume flux through BSO did not show a strong interannual variability.

• There was increase in volume flux in both the Fram Strait and over



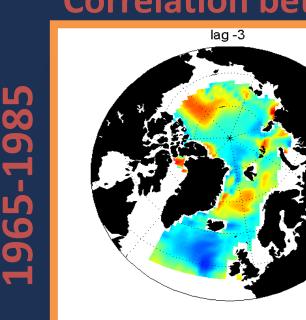
the GS Ridge for the period 1985-1995.

• The BSO and Fram Strait were the only straits that showed a monotonic increase in their decadal

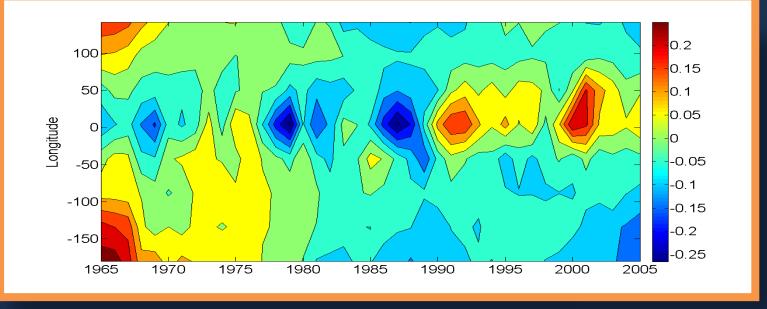
mean temperatures.

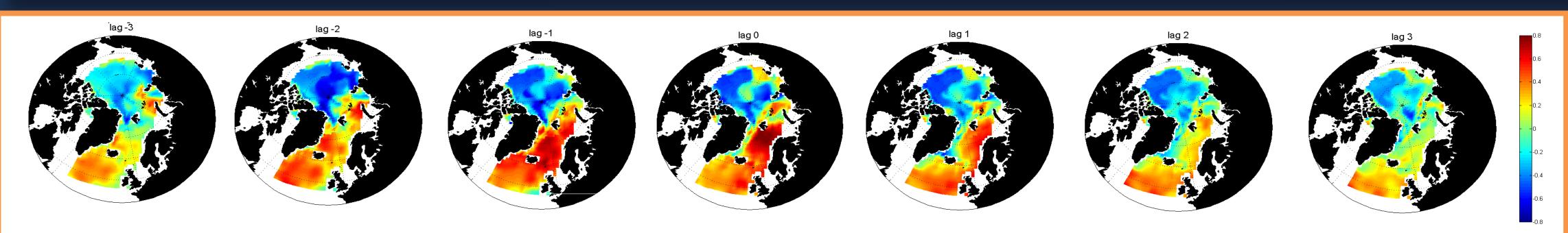
Possible Regime Shift In Heat Transport

• The bottom water core temperature (minimum temperature) close to the Fram Strait is anomalously warm after 1990 and close to the Bering Strait is anomalously cold after 1990 and vice versa before 1990.



Normalized Anomaly of Bottom water core temperature





• The correlation between maximum temperature of the water mass and maximum temperature in Fram Strait (lagged) changed between the two time periods. • Before 1985 there is only a strong correlation between the max temperature in the Fram Strait at zero and one year lags. After 1985 the correlation increases at lags larger than a year. • The propagation of anomalies of the AW temperature follow a counter-clockwise pattern the Arctic Ocean for the period 1986-2005. This pattern is not present in the period

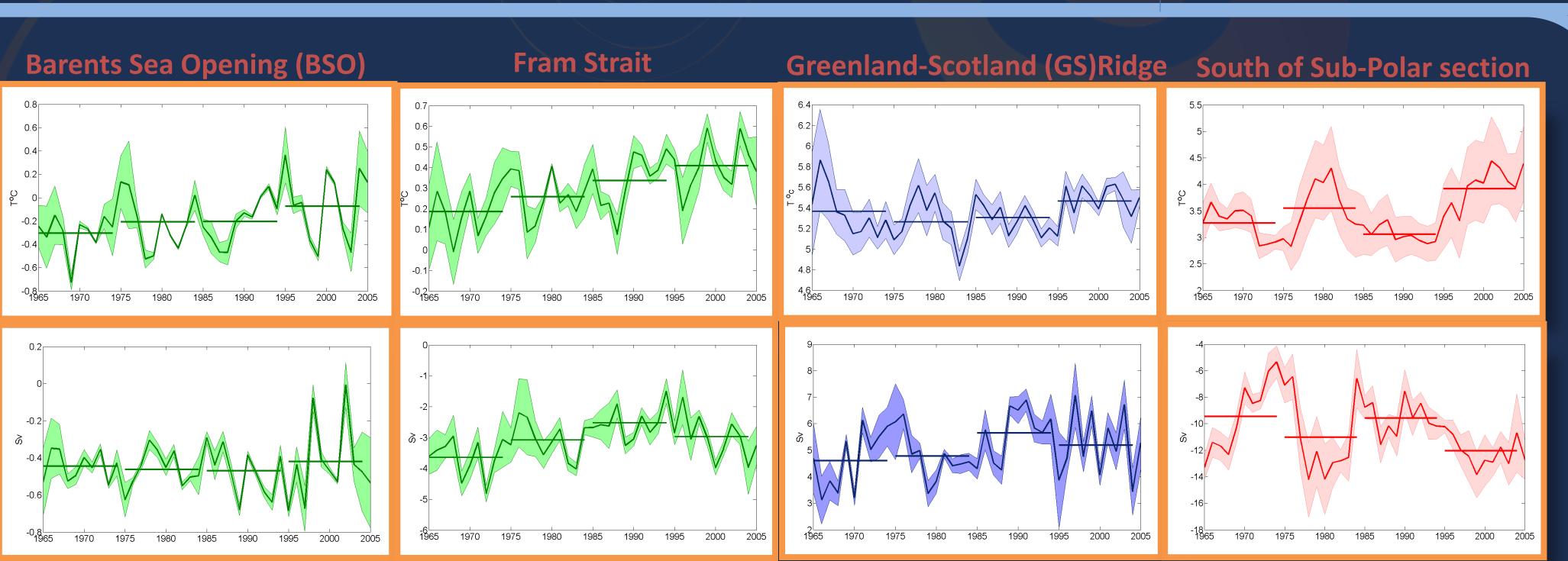
1965-1985.

Model Validation

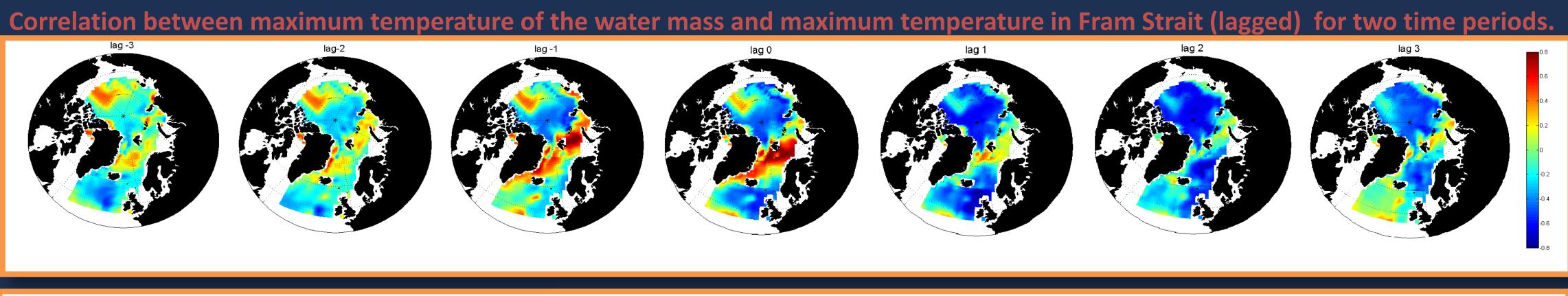
• Volume fluxes of Atlantic Water (AW) calculated from model results over the Greenland-Scotlan through the various branches shown in the table below, are within error of the volume fluxes of found using observations (Hansen 2008).

• The temperature of AW in all braches flowing over the Greenland-Scotland Ridge are colder in the

$ \begin{array}{c} $	a the model recults		Branch	Vol. Flux Sv	T °C	Sa
$\begin{array}{c c} \hline \\ \hline $		Observations 1999-2001	Iceland (I)	0.8 ± 1	6.0	≤ 3
$\begin{array}{c c} \hline \\ \hline $			Faroe (F)	3.8 ± 1	8.2	35
$\begin{array}{c c} \hline \\ \hline $			Shetland (S)	3.8 ± 1	9.5	35
Iceland (I) 1.0 ±0.1 5.3±0.3 3.9 Faroe (F) 4.0±1 6.7±0.3 3.9 Shetland (S) 5.2±1 8.3±0.2 3.9			Total Atl.	8.5 ±3	8.5	35
Parce (F) 4.0±1 6.7±0.3 3.1 Shetland (S) 5.2±1 8.3±0.2 3.1		1	Iceland (I)	1.0 ±0.1	5.3±0.3	35
$2 + \frac{1}{2} = $		del -200	Faroe (F)	4.0±1	6.7±0.3	35
	12°W 0°	оМ 9999.	Shetland (S)	5.2±1	8.3±0.2	35
Total Atl. 10 ± 2 6.8±0.3 3		H	Total Atl.	10 ± 2	6.8±0.3	35



• The decadal mean of the volume flux and temperature of the water entering into the south of the sub-polar section fluctuated between positive and negative anomalies. However a positive anomaly in temperature corresponds with a negative anomaly in volume flux and vice versa.



Fram Strait

 Northward and southward model volume flux through the Fram Strait are close to those calculated from observations (Schauer 2008). However there are smaller magnitudes in the model results than in observations.

nd	Ridge
fΑ	Ŵ

alinity

35.00

5.23

5.32

35.25

5.10±0.01 5.21±0.04 5.27±0.01

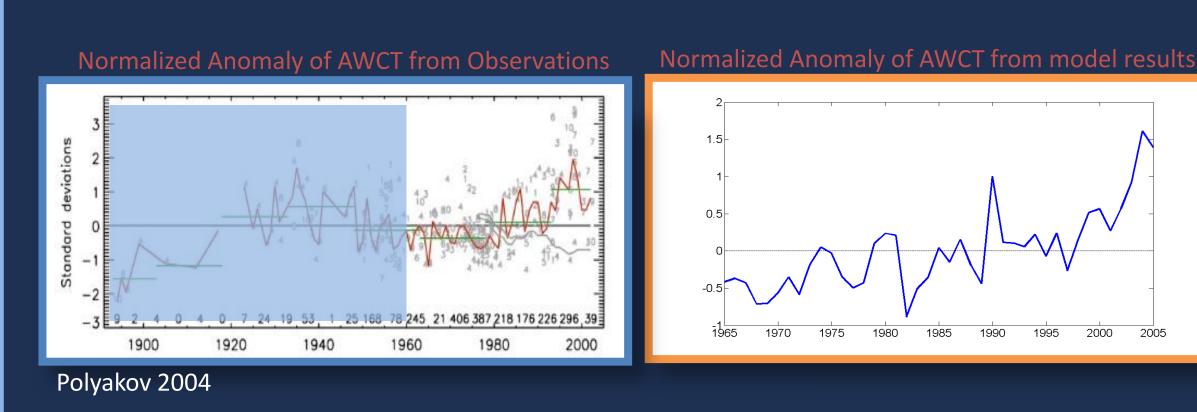
5.19±0.03

	Branch: Fram Strait	Vol. Flux Sv
Observations 1997-2003	Northward Flux	12 ± 1
	Southward Flux	14 ± 1
	Total	2 ± 1 South
Model 1997-2005	Northward Flux	7.5 ± 0.7
	Southward Flux	10.7 ± 0.6
	Total	3.1 ± 0.5 South

• The model reproduces correctly the seasonal time variability of the temperature in the Fram Strait. • The warm season in both the model and observations occurs in the late fall early winter. • The year 2003 was an anomalously warm year and in 2002 there was a westward cold intrusion in both model and observations.

Atlantic Water Core Temperature (AWCT) in the Arctic Ocean

• The trend in the AWCT in the Arctic Ocean produced by the model is close to that of the observations. The normalized anomaly of the AWCT is steady and negative between 1965 and 1990 and sharply increases after 1990.



Conclusions

-The model represents realistically the volume and temperature transport into the Nordic Seas over the Greenland-Scotland Ridge and through the Fram Strait. It also reproduces the trend in the AWCT in the late 1990s. -The decadal mean temperature through the Fram Strait and southern boundary of the sub-polar section increased between 1965 and 2005. The temperature in the rest of the straits remained relatively constant. In all straits except the BSO there was an increased volume flux for the period 1985-1995.

-Between the period 1965-1985 and 1986-2005 there was a change in the correlation between maximum temperature in the Fram Strait and in the ocean; with a stronger, lag correlation for the later period

-Before 1990 the BWCT was anomalously cold near where the Atlantic water enters and anomalously warm where the Pacific water enters. After 1990 this tendency switched.

eferences

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