



## **II.2.1 Potential Predictability of Current and Future Climates**

## **II.2.2 Prognostic predictability from ensembles of coupled model simulations**

**Bill Merryfield**  
***CCCma***



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# Predictability Studies

- Climate dynamics can be characterized as having two components:

‘Signal’: dynamics *deterministic*, potentially predictable

‘Noise’: dynamics *random*, unpredictable

- In *diagnostic* predictability studies the variability of the climate system is partitioned into these two components,  
→ *potentially predictable variance fraction* =  $\sigma^2_{\text{Pred}} / \sigma^2_{\text{Total}}$
- *Prognostic* predictability studies look at rates of divergence from neighboring initial conditions “ppvf”



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“ppvf”
- *Prognostic* predictability studies look at rates of divergence from neighboring initial conditions
- Functions of *time scale, region, climate variable...*

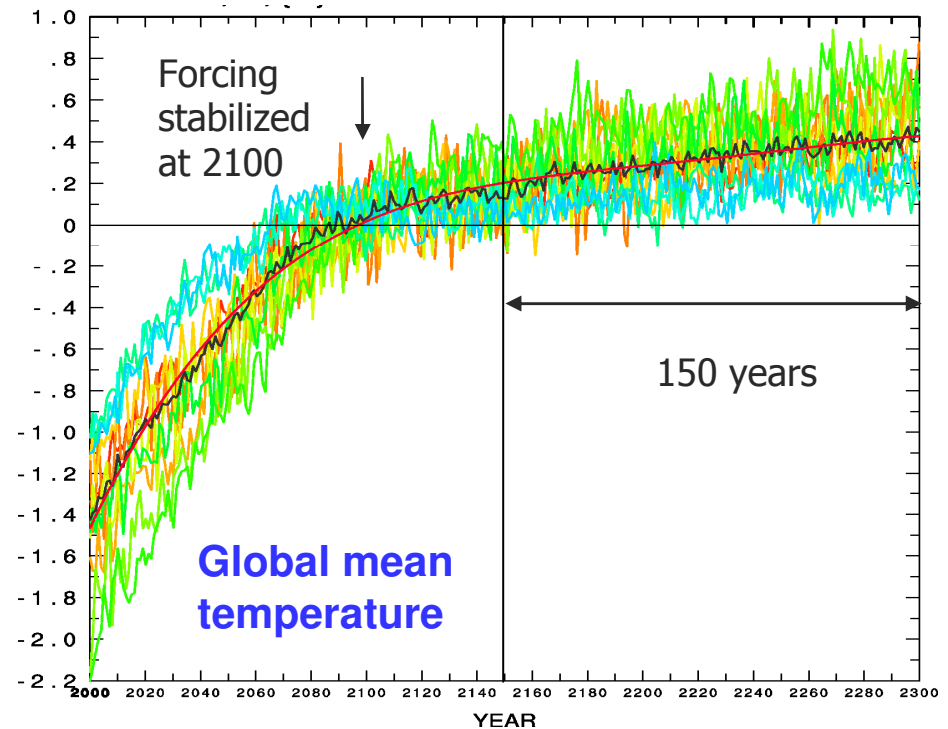


## II.2.1 Potential Predictability of Current and Future Climates

### Potential predictability in a warmer world

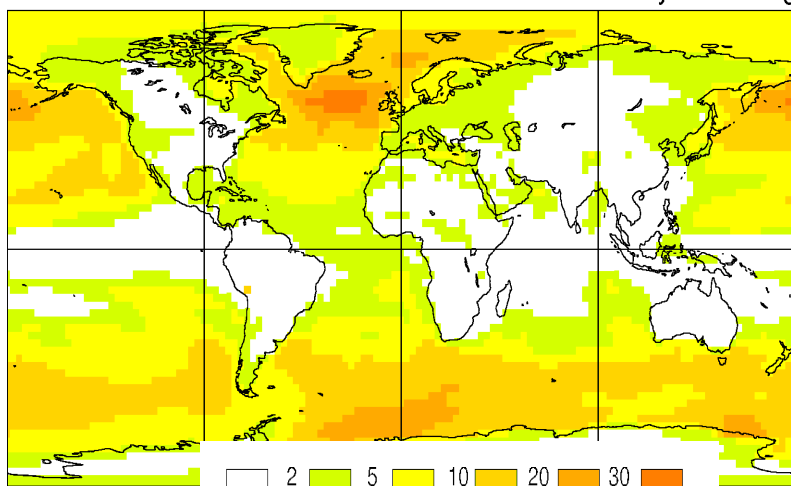
*Boer, J. Climate accepted*

- Consider B1/Stabilization Scenario
- Forcing stabilizes in 2100, consider 2150-2300
- 11 CMIP3 simulations
- Remove long term adjustments by fitting low-order polynomial

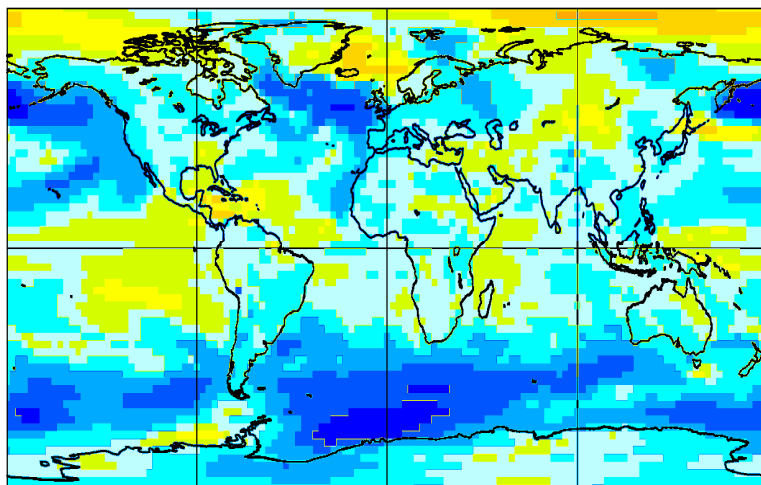
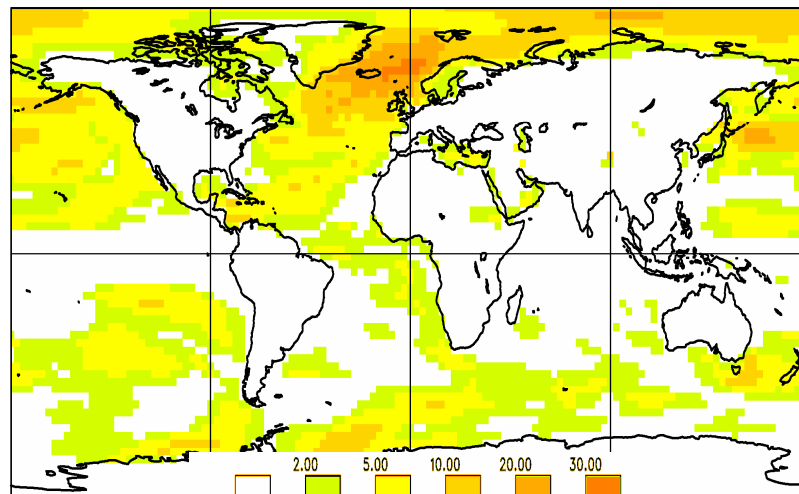


# Decadal ppvf (%) for Temperature

Control simulation

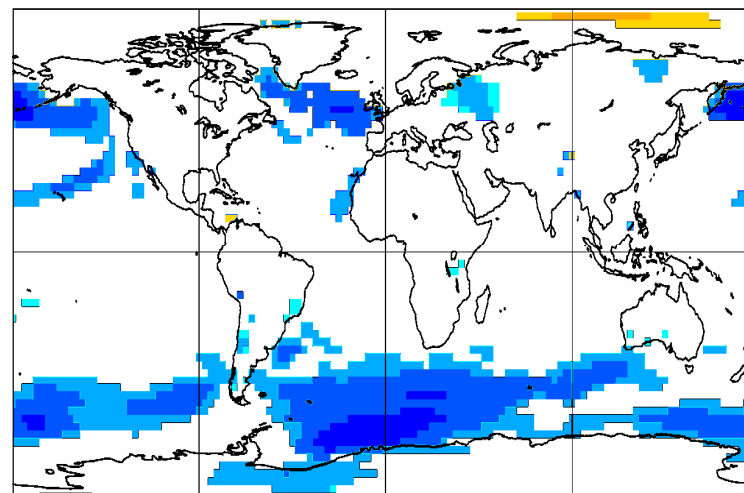


B1 stabilization



← less | more →  
 -20.00 -10.00 -5.00 -2.00 0.00 2.00 5.00 10.00 20.00

Difference in warmer world



-20.00 -10.00 -5.00 -2.00 0.00 2.00 5.00 10.00 20.00

Where confidence bands *don't* overlap



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## II.2.1 Potential Predictability of Current and Future Climates

### Decadal potential predictability of 21<sup>st</sup> Century climate

*Boer (2009) submitted*

- Decadal climate variations are superimposed on warming trend
- Decompose variance as

$$\sigma^2 = \sigma^2_{\Omega} + \sigma^2_v + \sigma^2_{\varepsilon}$$

externally forced component      internal variability component      "noise"

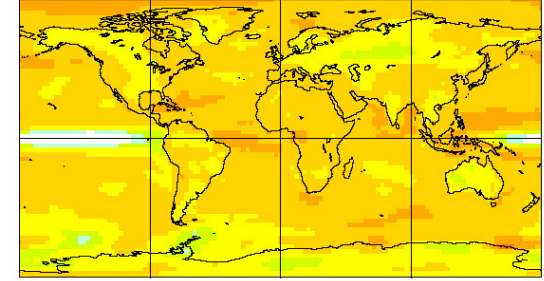
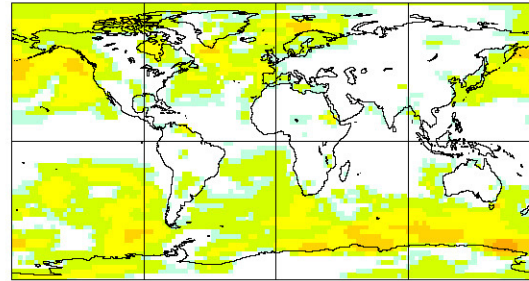
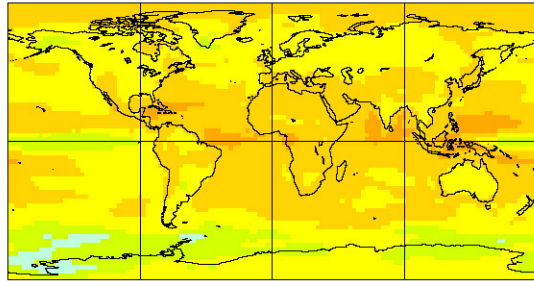
- B1 scenario 2001-2100
- CMIP3 → 35 simulations/18 models
- Examine decadal *forced* and *internal* ppvf

$$\sigma^2_{\Omega} / \sigma^2$$

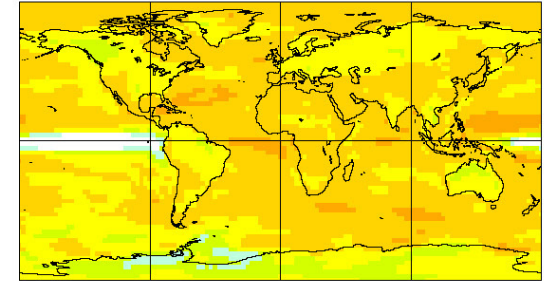
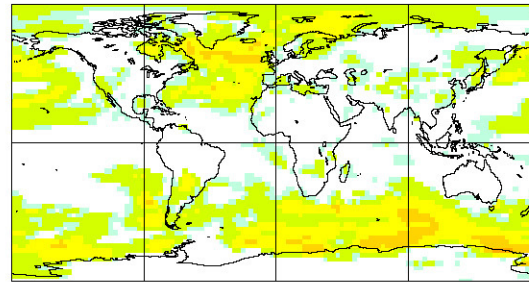
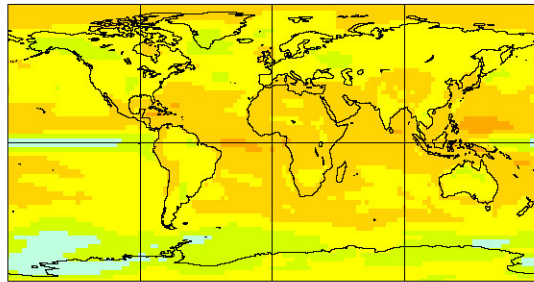
$$\sigma^2_v / \sigma^2$$



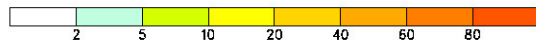
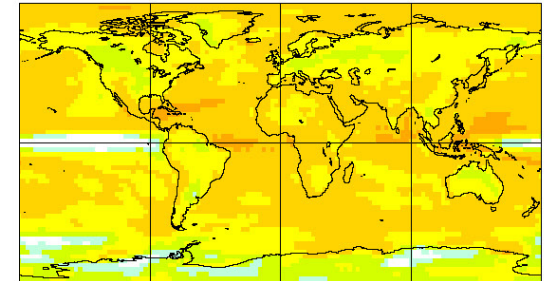
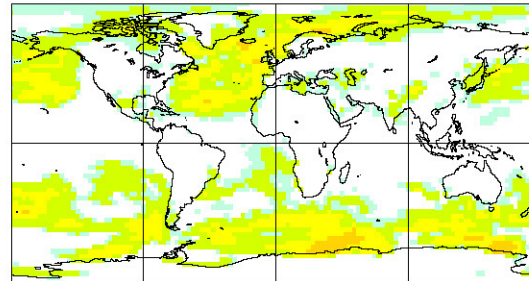
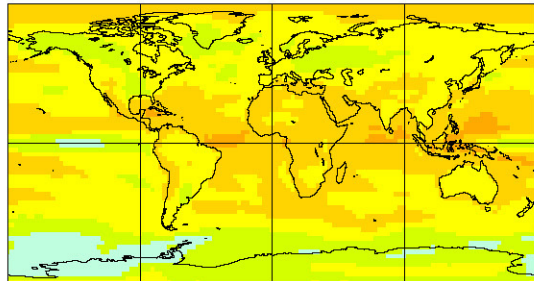
## Next-decade potential predictability 2020-2030



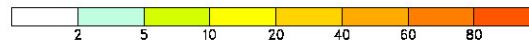
## 2030-2040



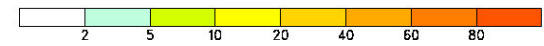
## 2040-2050



Forced component  $p_{\Omega}$



Internal component  $p_v$



Forced plus internal

→ these predictabilities should characterize decadal forecasts\*

\*to extent models are reliable



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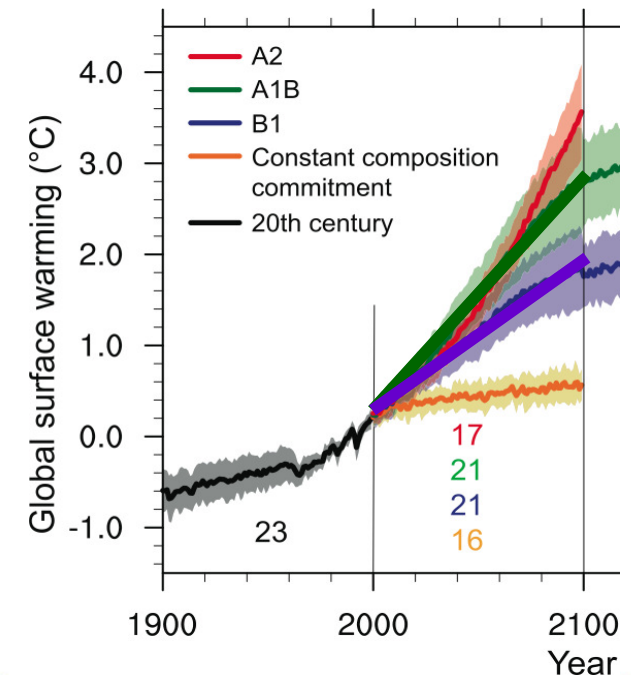
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## II.2.1 Potential Predictability of Current and Future Climates

### Likelihood and predictability of cooling episodes in a warming climate

*W. Merryfield and Ajayamohan Ravindran*

- Decadal forecasts from Kiel & Hadley Centre predict near-term *cooling*,
- *How likely is it that the next  $N$  years will be cooler than the last  $N$  years?* → probability  $P_N$
- Approach:
  - diagnose  $P_N$  directly from CMIP3 output\*:
    - B1 : 20 models, 48 runs, **4800 years**
    - A1B: 12 models, 12 runs, **1200 years**
  - apply statistical models



\*courtesy S. Lambert



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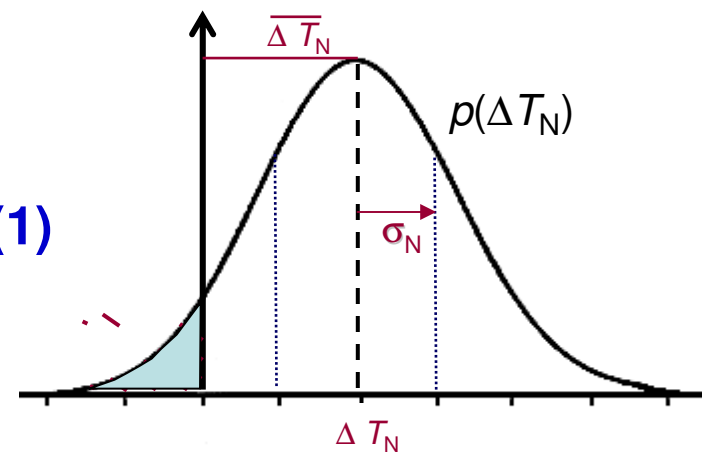


# Relate cooling probability to mean trend, internal variability

- Consider  $p(\Delta T_N)$  for *differences* between successive N-year means
- If  $p(\Delta T_N)$  is Gaussian then  $p \propto \exp - \frac{(\Delta T_N - \overline{\Delta T_N})^2}{2\sigma_N^2}$ 
  - mean trend
  - internal variability
- Cooling probability  $P_N$  corresponds to probability of  $\Delta T_N < 0$

$$P_N = \int_{-\infty}^0 p(\Delta T_N) d(\Delta T_N)$$

$$= \frac{1}{2} [1 + \text{erf}(-\overline{\Delta T_N} / \sigma_N \sqrt{2})] \quad (1)$$



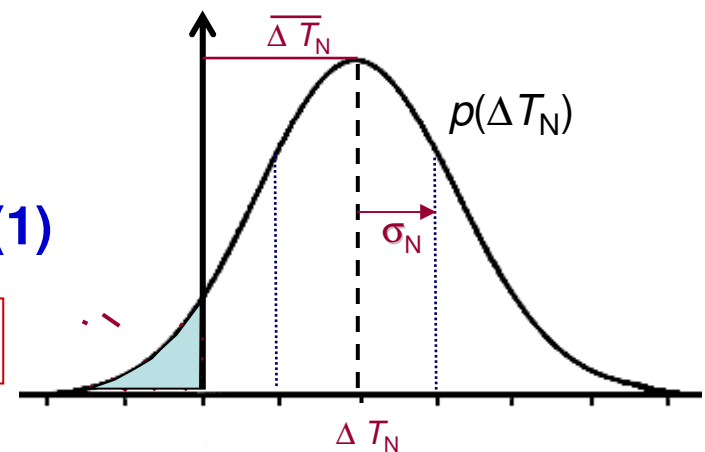
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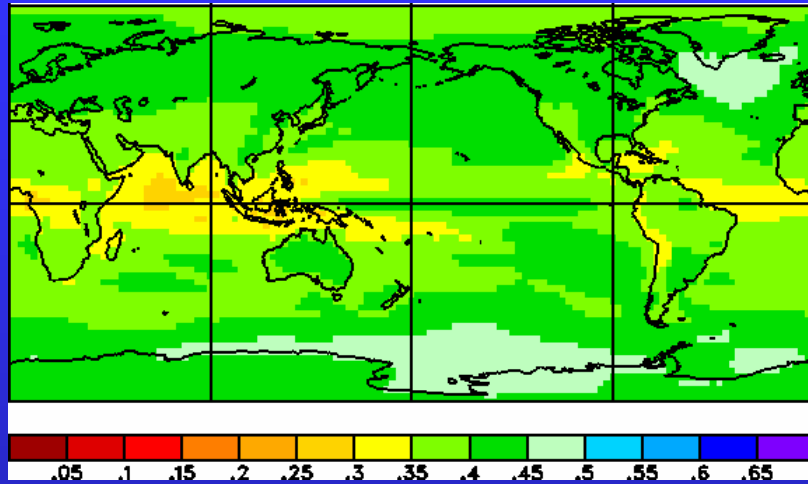
$$= \frac{1}{2} [1 + \text{erf}(-\overline{\Delta T_N} / \sigma_N \sqrt{2})] \quad (1)$$

→  **$P_N$  large where  $\Delta T_N$  small or  $\sigma_N$  large**



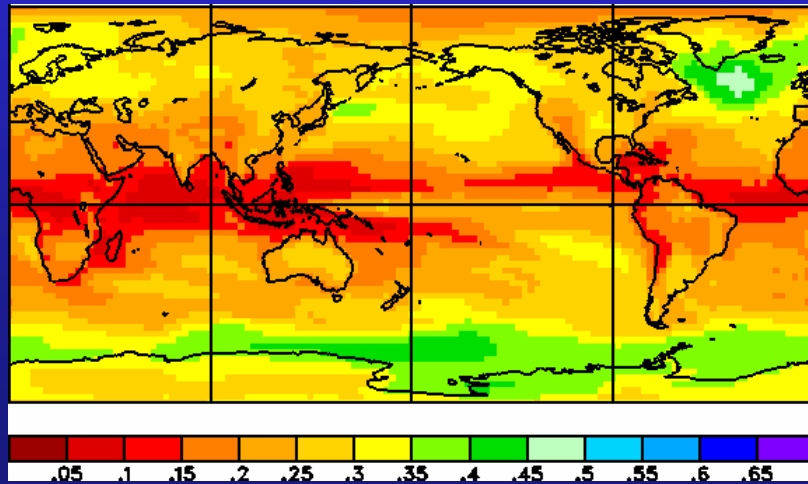
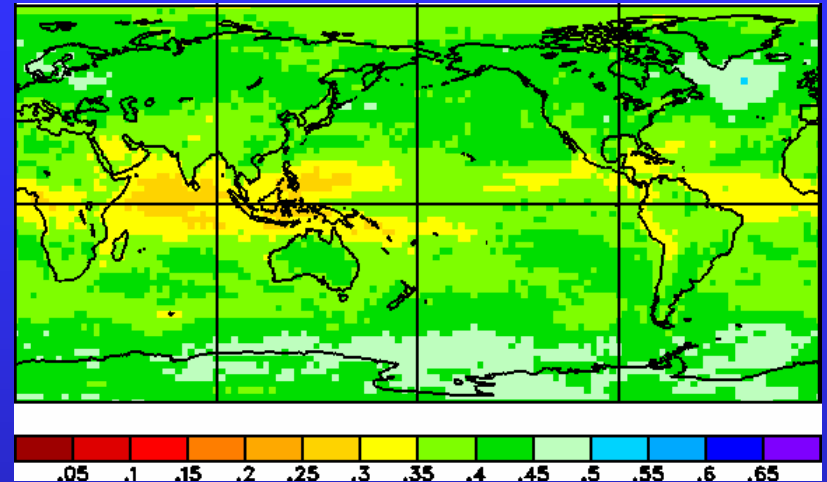
# Multi-model $P_N$ (B1)

From (1)

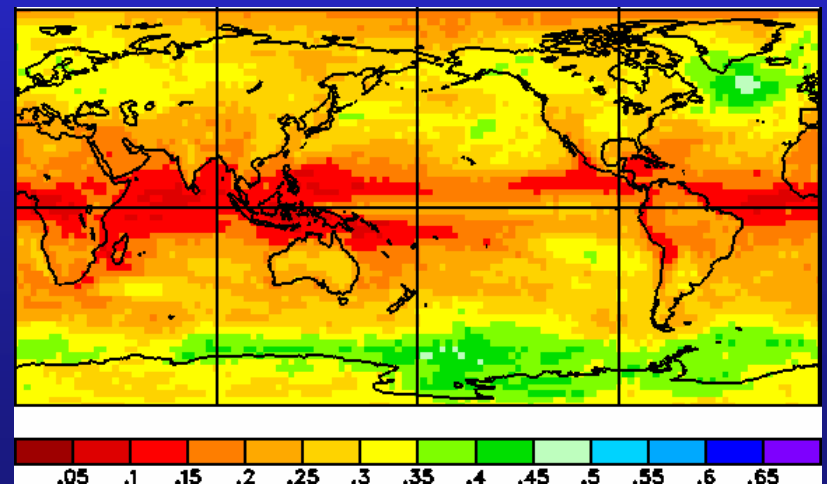


N=5

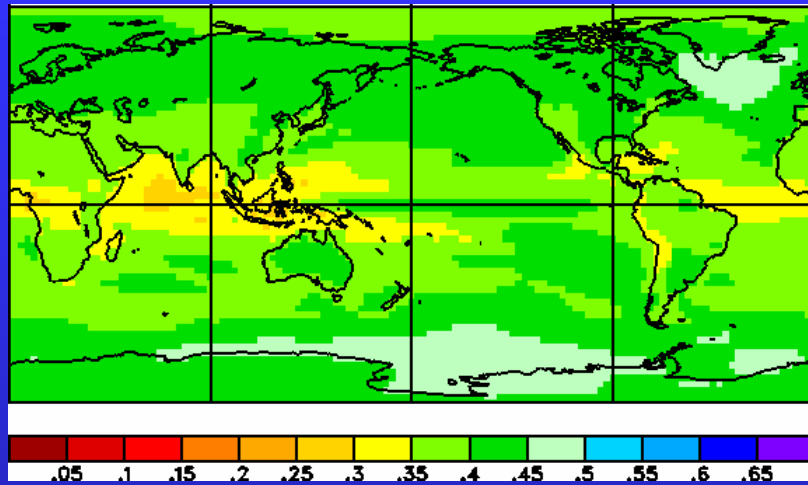
Direct calculation



N=10

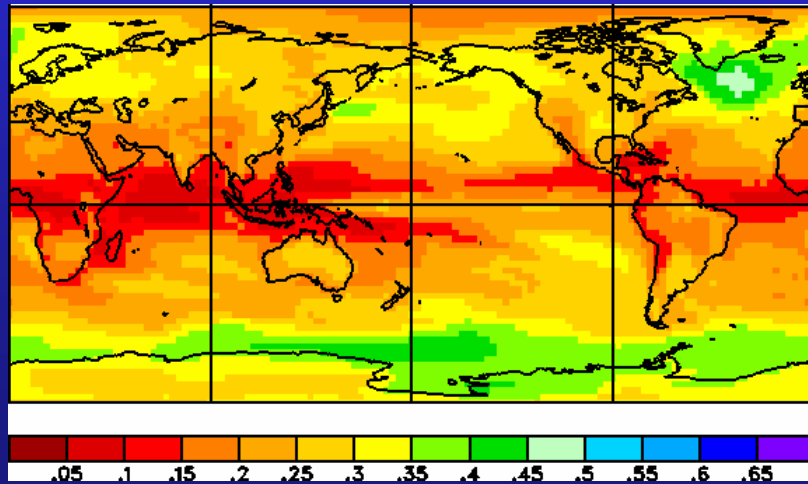
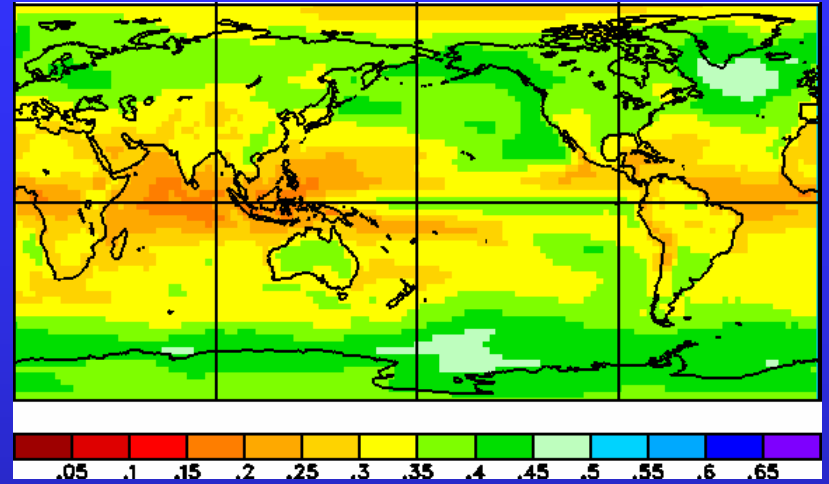


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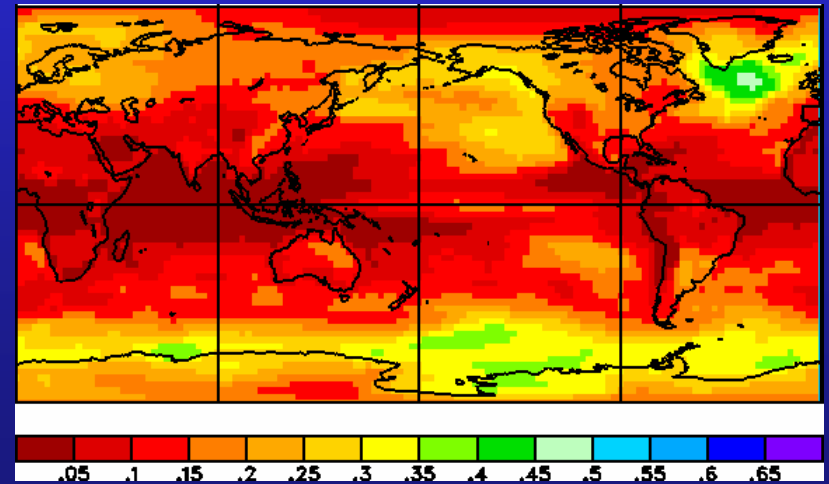


N=5

# Multi-model $P_N$ (A1B)



N=10



# Global mean T: Modeled vs Observed

	B1		$P_N$ from (1)
N	$\overline{\Delta T}_N$	$\sigma_N$	
5	0.081	0.081	15.8%
10	0.164	0.071	1.0%

	GISS		$P_N$ from (1)	HadCRUT3		$P_N$ from (1)
N	$\overline{\Delta T}_N$	$\sigma_N$		$\overline{\Delta T}_N$	$\sigma_N$	
5	0.079	0.046	4.3%	0.079	0.057	8.3%
10	0.159	0.011	0.0%	0.158	0.009	0.0%

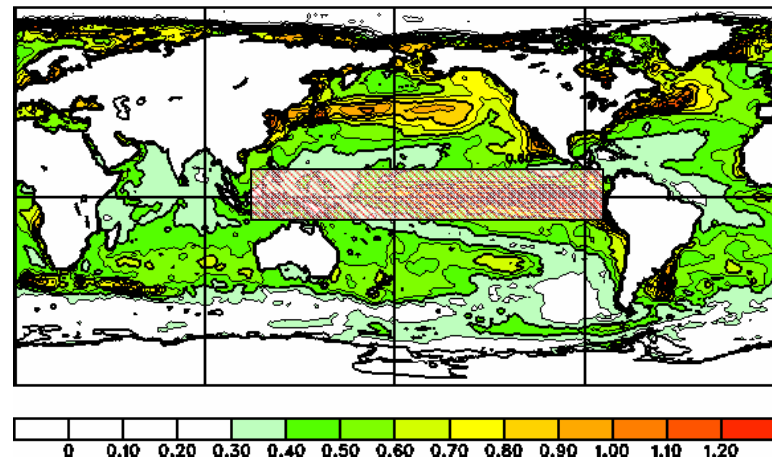
⇒ rate of warming similar to B1, but less variability (esp decadal)

## II.2.1 Potential Predictability of Current and Future Climates

### Regional impacts of air-sea coupling on climate variability and predictability

*Ajayamohan Ravindran, W. Merryfield,  
S. Kharin, G. Boer*

- Examine climate variability and potential predictability when atmosphere sees only *climatological* SSTs in specified regions:



- Ravindran afternoon talk...



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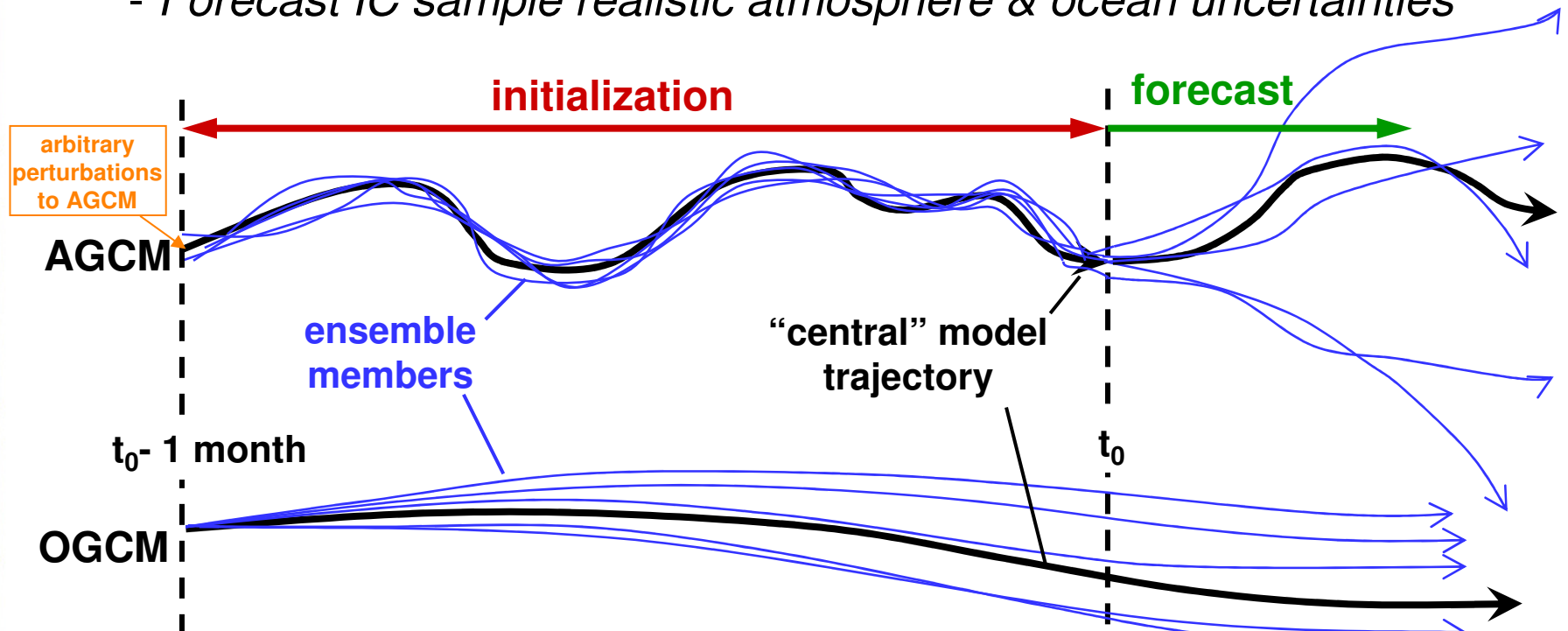
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## II.2.2 Prognostic Predictability from coupled model ensembles

- Aim: “perfect model” predictability experiment based on large (~100-member) ensemble of coupled model integrations
- Take advantage of new initialization technique: incremental analysis updates (IAU):
  - AGCM assimilates “central” model run for ~1 mon prior to fcst
  - *Forecast IC sample realistic atmosphere & ocean uncertainties*



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